Formation and Extension of Compaction Bands in Porous Sandstone

J. W. Rudnicki
Dept. of Civil and Env. Eng.
and
Dept. of Mechanical Eng.
Northwestern University
Collaborators

- Bill Olsson, Dave Holcomb (Sandia, Albuquerque)
- Eric Grueschow (Exxon-Mobil)
- Kathleen Issen (Clarkson University)
- Sheryl Tembe, Teng-Fong Wong (SUNY Stony Brook)
- Kurt Sternlof, Dave Pollard (Stanford)
- Jose Andrade (NU, now Caltech), Steve Sun (NU, now Sandia Livermore), Nicola Lenoir (NU, now Ecole des Ponts ParisTech)
- Peter Eichhubl (Texas Bureau of Economic Geology)

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Outline

• What are compaction bands?
• Why are they important?
• What do they look like?
• Why do they form?
• Models for propagation
What are compaction bands?

- **Compaction band**
- **Shear band (fault)**
- **Shear-enhanced compaction band**
- **55 – 65°**
- **10 – 20°**

Compaction bands are zones where materials are compacted due to shear stress. They are associated with certain angles of shear and can be enhanced by shear forces.
Why are compaction bands important?
From Vajdova, Baud and Wong, JGR, 2004
Permeability evolution during localized deformation in Bentheim Sandstone

In situ permeability measurements inside compaction bands using X-ray CT and lattice Boltzmann calculations
Lenoir, Andrade, Sun and Rudnicki, Proceedings of GEOX 10, 2009
Shortest Flow Path Inside and Outside Compaction Bands

INSIDE CB
ϕ = 0.14

τ = 2.79
K = 3.4e-13 m²

τ = 2.15
K = 5.3e-13 m²

τ = 2.56
K = 4.4e-13 m²

OUTSIDE CB
ϕ = 0.21

τ = 1.77
K = 1.3e-12 m²

τ = 1.76
K = 1.2e-12 m²

τ = 1.81
K = 1.3e-12 m²
Uniform compaction

Localized compaction
What do compaction bands look like in the field?
Figure 1. Location of the Valley of Fire State Park, southeastern Nevada (inset), where more than 20 square kilometers of the 1,400-m-thick eolian Jurassic Aztec Sandstone are extensively exposed. Photo shows view northward from the location marked * on the park detail. Throughout the upper half of the Aztec, compaction bands crop out in positive relief as sub-parallel, centimeter-thick, north-northwest-trending, steeply east-dipping tabular fins spaced from centimeters to meters apart.

Compaction Bands: Earliest Structural Fabric of the Aztec

From Kurt Sternlof, Stanford

Figure 3. Close-up of a typical, well-developed compaction band fin in outcrop. Note that depositional bedding extends relatively undisturbed across the band, and is clearly visible on the fin.
Figure 2. Left-hand photo shows an outcrop of widely spaced, relatively planar and parallel compaction bands along the northeast flank of Silica Dome. Arrows indicate opposite tips of a single band 62 m long and up to 15 mm thick (illusory gaps in band continuity are due to outcrop topography and breaks in the telltale fin). A total of 16 tip-to-tip thickness profiles were measured using a steel tape and calipers. Right-hand photo illustrates that, even when closely spaced, compaction bands in this locale tend to remain planar (arrow indicates tip).

From Sternlof, Rudnicki and Pollard, JGR, 2005
Figure 5. Photomicrograph of compaction band sampled 6.0 meters from the tip, where it is about 9 mm thick (dotted black lines). Blue indicates epoxy-filled pore space (~25% outside the band, ~10%
Fig. 4. Pure (PCB) and shear-enhanced (SCB) compaction bands exposed on a bedding plane. a. Pure compaction bands with characteristic wavy trace. Pocket knife for b. Chevron pattern of alternating right- and left-lateral shear-enhanced compaction bands. Chevron-style bands merge with and continue as planar shear-enhanced compaction bands. c. Intersection of sets of shear-enhanced compaction bands. d. Chevron-type compaction bands with out-of-phase alignment of crests and troughs.
Compactive Shear Band with 2.5 cm reverse slip

Shear band with 1 cm slip
Macroscopic band is composed of multiple parallel bands of more intense grain-size reduction

Detail of

Eichhubl, Hooker, Laubach, J. Struc. Geology, 2010
shear enhanced compaction bands

Fossen, Schultz & Torabi
J. Structural Geology
in press

shear enhanced compaction bands

pressure solution

slipped compactional shear band
offset equal a few cm
What do compaction bands look like in the laboratory?
Bentheim sandstone: $P_c = 300 \text{ MPa}$

$\varepsilon_{ax} = 1.4 \%$

$\varepsilon_{ax} = 3.1 \%$

$\varepsilon_{ax} = 4 \%$

$\varepsilon_{ax} = 6 \%$

Baud, Klein, and Wong, Compaction localization in porous sandstones:…, JSG, 2004 GRL, 2001
A compaction band in the Bentheim sandstone: typically the lateral width extends over 2 grains or so (~ 600 μm)
Baud, Klein and Wong, Compaction localization in porous sandstones: spatial evolution of damage and acoustic emission activity, J. Struct. Geology, 2004

Why do compaction bands form?
Homogeneous and homogeneous deformation

Continued Homogeneous deformation

Shear band or compaction band

Stress

Strain
Kinematic Condition

\[ \tan \psi = \frac{\Delta D^\text{vol}}{\Delta D^\text{shear}} \]

\[ \mathbf{v}^\text{band} = \mathbf{v}^0 + \mathbf{g}(\mathbf{n} \cdot \mathbf{x}) \]

\[ \mathbf{D}^\text{band} = \mathbf{D}^0 + \frac{1}{2} (\mathbf{n} \mathbf{g} + \mathbf{g} \mathbf{n}) \]

For compaction band, \( \mathbf{g} = \alpha \mathbf{n} \),

\[ \psi = -90^\circ \]
Equilibrium Condition

\[ \frac{dt}{d\sigma} = n \cdot d\sigma_{\text{band}} = n \cdot d\sigma_{\text{o}} \]
Elastic-plastic rate-independent constitutive model depending on first and second invariants (Drucker – Prager type).

\[ d\varepsilon^p = \beta d\gamma^p > 0 \text{ (dilation)} \]

\((\sigma_a - \sigma_c) / \sqrt{3}\)
for standard test

\((\sigma_a + 2\sigma_c) / 3\)
for standard test

\(k \leq k_{\text{crit}}\) for localization
\( \theta_{\text{crit}} \) = angle between fault normal and most compressive principal stress.
\[ \tan \psi = \frac{\Delta \varepsilon}{\Delta \gamma} \]

Range of angles for shear–enhanced compaction bands by Eichhubl et al. (2010)

\[ \beta < 0 < \mu \]
compaction bands
no shear

No volume change

Range of $\beta$ and $\mu$ from data of Baud, Vajdova and Wong (JGR, 2006) for Adamswiller, Bentheim, Berea and Darley Dale sandstones. Typically $\mu > 0$ and $\beta < 0$ and small.
Evaluate for an Elliptic Yield Cap

Rudnicki, JGR, 2004

More elaborate yield surface model
Grueschow and Rudnicki, Int. J. Solids Struc., 2005
COMPACTIVE YIELD STRESSES OF 4 SANDSTONES:
The critical stress levels $C^*$ at the onset of shear-enhanced compaction were fitted with elliptical caps (yield envelopes).
\[
\frac{(\sigma - c)^2}{a^2} + \frac{\tau^2}{b^2} = 1
\]

\(\varepsilon = (a / b)^2 = 1.53\) to 3.0

from Wong, David & Zhu (JGR, 1997)

\(\varepsilon = 1.268\) for Bentheim Sandstone

from Klein et al. (Phys. Chem. Earth A, 2001)

axisymmetric compression path
Increasing $k_{crit}$

Shear bands

Elliptical cap surface

Increasing $k_{crit}$

$k_{crit} = 0$

(for normality)

Increasing confining stress

Increasing $k_{crit}$

Compaction bands
Figure 3. Transmission optical micrographs of deformed wet Diemelstadt sandstone samples. The effective pressure and the level of axial strain are indicated next to the thin sections. The width of each thin section is ~18 mm. Principal stress $\sigma_1$ was along the axial direction. Brittle failure in the Diemelstadt sandstone occurs at effective pressures of 10–60 MPa as (a) dilatant conjugate shear-bands, (b) compactant conjugate shear-bands, (c) diffuse conjugate shear-bands, and (d) mixed failure modes of high-angle shear-bands and compaction bands. At higher pressure (>60 MPa), discrete compaction bands are the dominant failure mode. The development of an array of discrete compaction bands at 150 MPa effective pressure can be traced with increasing axial strain: (e) 1.6%, (f) 2.1%, and (g) 4%. The black arrow in Figure 3g points to a layer of oxides separating coarser material below from finer material above. (h) Mosaic showing a discrete compaction band propagating across the sample in the area delimited by a rectangle in sample D13 (f). Significant comminution makes the band appear dark.
Open circles, peak stress at brittle failure
Solid circles, onset of shear – enhanced compaction

Fig. 15. Peak stresses and yield stresses at the onset of shear-enhanced compaction for Darley Dale (a), Berea (b), Rothbach (c) and Bentheim sandstones (d).
How do the bands propagate (extend)?
Figure 1: Compaction Band Data

Linear fit: $y = A + Bx$

$A = 0.43240, B = 0.49891$

after Tembe, PhD Thesis, SUNY, Stony Brooke
Ellipsoidal band with uniform inelastic compactive strain $\varepsilon^p$

compactive displacement $= \frac{w\varepsilon^p}{1 - \varepsilon^p} \approx w\varepsilon^p$

$2w = \text{band width}$

For aspect ratios typical of the field, $10^{-3}$ to $10^{-4}$, initial thickness can be neglected entirely.

“anti-crack” model
Fletcher and Pollard Geology, 1981
Combined “anti-crack” and “anti-dislocation” model

uniform compactive displacement, $2w$

uniform resistive normal traction

$a$ $a$

$L$ $L$
\[ K = \frac{\mu w}{(1 - \nu)} \sqrt{\frac{\pi}{L-a}} \left[ \frac{k \sqrt{L-a}}{E(k) - (1 - k^2)K(k)} \right] \]

where \( k^2 = 1 - \left(\frac{a}{L}\right)^2 \)

\[ G_{\text{crit}} = \frac{(1 - \nu)}{2\mu} K_{\text{crit}}^2 \implies K_{\text{crit}} = \sqrt{\frac{2\mu G_{\text{crit}}}{1 - \nu}} \]

\[ 2w = \sqrt{L} \sqrt{\frac{8(1 - \nu)G_{\text{crit}}}{\pi \mu}} F(k), \quad L \gg a \approx 1 \]

For \( \nu=0.2, \mu = 8.333 \text{ GPa}, G_{\text{crit}} = 40 \text{ kJ/m}^2 \)

\[ \implies G_{\text{crit}} / \mu = 4.8 \times 10^{-3} \text{ mm} \]

(Sternlof, Rudnicki & Pollard, 2005)

\[ \frac{2w}{\sqrt{L}} = 0.003127 \text{ m}^{1/2} \]
Compaction Band Data

Linear Fit: $y = A + Bx$

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<th>$A$</th>
<th>$B$</th>
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Conclusions

• Compaction bands are predicted to occur on portions of the cap surface where they are (roughly) observed to occur.
• Predictions are roughly consistent with lab data but do not agree very well quantitatively.
• Need for better constitutive data and modeling and a better understanding of the micromechanical factors affecting macroscopic response.
• More complex load paths (now being used) should be more useful for constraining constitutive models.
Conclusions

• A simple fracture and dislocation model indicates a (surprising) consistency between lab and field data.
• Model is, however, speculative since little information is available on the initiation and propagation of bands in the field.
• Field and laboratory observations of the breakdown process near the tip of bands would be helpful in constructing more elaborate models.
Questions

• How do lab, field and theoretical results fit together?
• What properties (or factors) other than porosity affect compaction band formation?
• What is a good physical microstructural microstructural model?
• How do microscale properties and lithology translate to macroscopic material (constitutive) behavior?
• What controls band thickness (and changes in thickness) and spacing?
• If bands formed under saturated conditions how does interaction of fluid flow with deformation affect formation and extension?
• How do we model the transition from initial band formation to extension of a fully formed band?
Thanks!

Questions for me?